

Fast secondary aerosol formation in southern African biomass burning smoke

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Biomass burning (BB) is one of the largest sources of reactive trace gases and fine particles in the Earth's atmosphere, significantly affecting the global radiation balance (e.g. Bond *et al.*, 2013). However, there is large uncertainty in the climate effect of the BB emissions, mostly because of the reactive species co-emitted with black carbon (BC) particles (Bond *et al.*, 2013).

We investigated the temporal evolution of BB smoke in southern Africa utilising continuous measurements at Welgegund, South Africa (26.57°S, 26.94°E, 1480 m above sea level) from 20 May 2010 to 31 December 2014. Of the measurements performed at Welgegund (welgegund.org), we used aerosol particle number size distribution, carbon monoxide (CO), BC with a MAAP, local wind and solar radiation. Submicron particulate mass (PM₁) was estimated from the number size distribution using a constant density of 1.5 g cm⁻³.

Using the methodology of Vakkari *et al.* (2014) we could identify 118 BB plumes over the nearly five year period. For 83 plumes the location of the fire could be identified from MODIS burnt area and/or SEVIRI fire radiative power observations. For these episodes – a total of 250 hours of in-plume sampling – we estimated the plume age based on local wind speed and direction (c.f. Vakkari *et al.*, 2014). For each plume the excess CO, BC and PM₁ (denoted with a Δ) were calculated by subtracting respective background concentrations, with Δ CO, Δ BC and Δ PM₁ given as $\mu\text{g m}^{-3}$ at STP.

In order to estimate the ratio of flaming to smouldering combustion, we calculated the Δ BC/ Δ CO ratio for each 10-minute Δ PM₁ sample: low Δ BC/ Δ CO indicates smouldering combustion and high Δ BC/ Δ CO designates flaming combustion. According to plume transport time we could also estimate the time each 10-minute sample had been transported in daylight, i.e. when global radiation > 100 W m⁻² at Welgegund. Combination of these parameters allowed us to investigate the effects of burning conditions and atmospheric transport on PM₁ mass in BB plumes, both in daylight and in the dark.

Smouldering BB smoke (Figure 1a) exhibits fast secondary aerosol formation in daylight: up to a factor of 4 increase in Δ PM₁/ Δ CO in less than 5 hours. Previously, aged BB plumes at Welgegund have been shown to be dominated by organic aerosol (OA) (Vakkari *et al.*, 2014), suggesting that the increase in Δ PM₁/ Δ CO is mostly due to formation of secondary organic aerosol (SOA). As samples with higher and higher Δ BC/ Δ CO are considered, the increase in Δ PM₁/ Δ CO becomes slower and even decrease for the most flaming cases (Figure 1b).

In the dark both increase and decrease rates in Δ PM₁/ Δ CO are an order of magnitude slower than in daylight, suggesting that the rapid daytime changes are driven by photochemical reactions: SOA formation in smouldering BB smoke and fragmentation of primary OA in flaming BB smoke.

To conclude, we have shown that during daytime secondary aerosol formation can increase the PM₁ mass in BB smoke by up to a factor of 4 in a few hours, but the increase depends critically on whether the fire was flaming or smouldering.

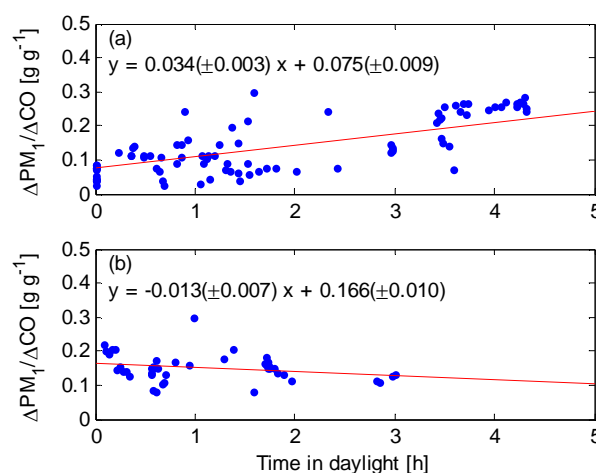


Figure 1. Daytime Δ PM₁/ Δ CO time evolution for (a) smouldering cases, Δ BC/ Δ CO < 0.01 g g⁻¹ and (b) flaming cases, Δ BC/ Δ CO > 0.03 g g⁻¹. A bivariate linear fit with standard error is included in both (a) and (b).

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